



Article

Chemical-Petrographic Types and Shock Metamorphism of 184 Grove Mountains Equilibrated Ordinary Chondrites

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Abstract: We reported the petrography and mineral chemistry of 184 equilibrated ordinary chondrites collected from Grove Mountains, Antarctica. The chemical-petrographic types and shock metamorphism degrees of these chondrites were assigned. They were classified into 46 H groups (22 H4, 20 H5, and four H6), 133 L groups (eight L4, 75 L5, and 50 L6), and five LL groups (four LL4 and one LL5). Some of these chondrites could be paired; however, both H and L group meteorites were affected. Further studies such as terrestrial ages and thermal luminescence are required in order to confirm the pairings. The relative abundances of H, L, and LL are different in Grove Mountain meteorites, when compared to those in Transcontinental Ridge meteorites. Based on the shock effects, the shock metamorphism degrees of these chondrites were assigned. Compared to previous studies, the heavily shocked samples of S4 and S5 have a higher fraction (59 out of 184) in Grove Mountain ordinary chondrites. The L group (54 out of 59) is the dominant chemical group in the heavily shocked chondrites, except for five meteorites which belong to the H group. The shock metamorphism degrees of the H and L groups are distinct, which may indicate different surface properties in their parent bodies. In addition, the petrologic types and shock degrees are probably closely related, with the most heavily shocked chondrites observed in types 5 and 6.

Keywords: chemical-petrographic types; shock metamorphism; chondrites; Grove Mountains; Antarctica

1. Introduction

Since the first discovery of nine meteorites of different chemical groups in the Yamato area, Antarctica by a Japanese team in 1969, Antarctica has become the most meteorite-rich region in the world. Many rare types of meteorites have been found in Antarctica, e.g., Martian meteorites, Lunar meteorites, and carbonaceous chondrites. The Grove Mountains consist of 64 nunataks, which are located in eastern Antarctica [1]. During the 15th Chinese Antarctic exploration between 1998 and 1999, a Chinese team conducted their first field survey of the Grove Mountains. More than 10,000 meteorites were collected by the Chinese Antarctic Research Expedition (CHINARE) in this area (Figure 1).

With a large number of Antarctica meteorites found, some comparative studies could be conducted, e.g., the type and mass distribution patterns and shock metamorphism in different chemical and petrographic chondrite types. Previously, some studies showed that there is no close connection between the distribution patterns of the types and mass, and the sites where the meteorites are collected [2–4]. On the other hand, some large meteorite showers have been evidenced in Antarctica. Locally, these showers do result in establishing a connection between meteorite types,

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mass, and sites [5]. Shock metamorphism from hypervelocity collisions is the most common feature of meteorite parent bodies. Deformation, melting, and the decomposition of mineral components are recorded under shock effects [6,7]. Moreover, natural high-pressure polymorphs of rock-forming minerals were also discovered in heavily shocked meteorites [6–10]. Since the high pressures and temperature conditions during shock events are comparable to those of the Earth's transition zone or of the lower mantle, the study of shock effects in meteorites is of great importance for deciphering the collisional and geological history of their asteroidal parent bodies [11,12]. Based on the shock effects in silicate minerals and shock-induced localized melting in ordinary chondrites, seven degrees of shock metamorphism (S1–S6 and shock melted) were classified by Stőffler et al.; the Pressure-Temperature (P-T) conditions of every shock degree were also given in their study [13].

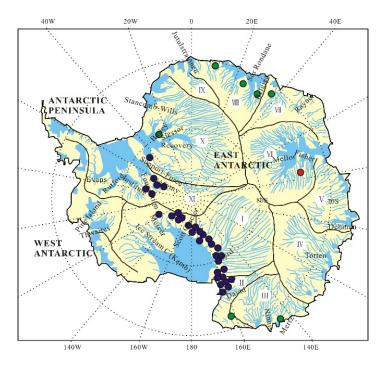


Figure 1. Map of meteorite concentration sites in Antarctica (based on the literature [14]). The red dots, blue dots, and green dots represent the meteorite enrichment areas found by China, the United States, and Japan and other countries, respectively.

In order to compare the type distribution patterns and shock metamorphism degrees with different chemical-petrographic types in chondrites, 184 equilibrated ordinary chondrites were randomly selected for classification. In this paper, the petrography and mineral chemistry of these meteorites are reported, and their chemical-petrographic types and shock metamorphism degrees are assigned.

2. Samples and Experiments

Samples of these meteorites were embedded in epoxy, and then cut into <1 mm thin slices with a low-speed diamond saw. They were prepared to standard polished thin sections without water. Textural observations of these chondrites were studied using an optical microscope and a back-scattered electron (BSE) image model of a JEOL 8100 electronic probe microanalyzer (EPMA) (Tokyo, Japan) at the Guangzhou Institute of Geochemistry, Chinese Academy of Sciences. Quantitative analyses for silicates were carried out using the same EPMA with a 15 keV accelerating voltage and a 20 nA beam current. Natural and synthetic minerals were used as standards. An X-ray peak deconvolution program was applied to correct for the K_{α} lines of V and Mn by the K_{β} lines of Ti and Cr. Analytical results of the silicates ware corrected by the ZAF method. Modal contents of opaque mineral (metallic Fe-Ni and troilite) were calculated from areas of the phase in BSE images of the polished thin sections.

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Chemical-petrographic types of these chondrites are mainly based on References [15–17], and their shock metamorphism results are based on Reference [13].

3. Chemical-Petrographic Types

The basic characteristics of these Grove Mountains chondrites are summarized in Table A1, including chemical-petrographic types, found location, mass, opaque mineral (vol %), mineral chemistry, and shock metamorphism. They are all equilibrated, including 46 meteorites classified as H chondrites (22 H4, 20 H5, and four H6), 133 L chondrites (eight L4, 75 L5, and 50 L6), and five LL chondrites (four LL4 and one LL5).

All 184 meteorites have homogenized mineral chemistry and their matrices are recrystallized (Figure 2). This indicates that they experienced significant thermal metamorphism in their parent bodies. Type 5 and 6 chondrites have blurred outlines of chondrules (Figure 2c,d), but chondrules are clearly outlined in type 4 (Figure 2a,b). The PMD (percent mean deviation) value of the Fa (Fayalite) content of olivine in all chondrites is most <5%. Low-Ca pyroxene is also homogeneous in most of these chondrites, with a PMD of Fs (ferrosilite) content <5%. The Fa content of olivine and Fs content of low-Ca pyroxene of these meteorites are plotted in Figure 3.

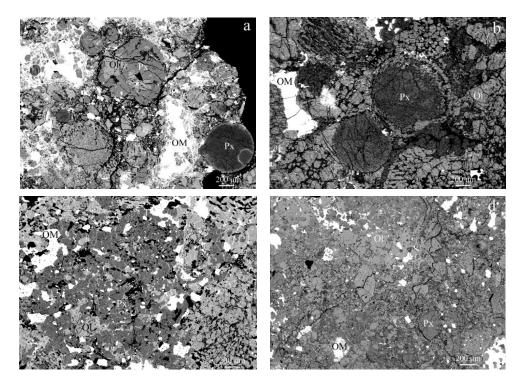


Figure 2. Back-scattered electron (BSE) images of representative equilibrated ordinary chondrites. (a) GRV 020008 (H4); (b) GRV 020038 (L4); (c) GRV 022178 (L6); (d) GRV 020041 (LL4). Ol—olivine, Px—pyroxene, OM—opaque mineral.

3.1. Group H

There are 46 meteorites in this group (Table A1). The Fa values of olivine for all chondrites range from 17.0 to 21.0 mol %, and the Fs values of low-Ca pyroxene range from 15.4 to 18.5 mol %. Both the Fa and Fs tally with the ranges of the H group (Figure 3). These 46 meteorites experienced significant thermal metamorphism. Twenty-two meteorites (Table A1) have a homogenized mineral chemistry and clear outlines of chondrules. Additionally, their fine-grained matrices are partly recrystallized (Figure 2a), with the sizes of the recrystallized matrices being $5-15~\mu m$. Brown-colored glass was found in a few chondrules. Accordingly, these meteorites belong to H4. The outlines of the chondrules in 20 chondrites (Table A1) can been recognized. Additionally, their matrices are well

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recrystallized, with the sizes of these grains being 10– $40~\mu m$. Secondary feldspar within the matrix and chondrules is popular, with sizes ranging from 10– $50~\mu m$. These 20 chondrites were assigned to H5. They were identified as GRV 020072, 020093, 020098, and 021522, having similar petrographic characteristics. Additionally, their matrices are intensely recrystallized, with the sizes of these grains being 50– $150~\mu m$. The chondrules in these meteorites are hard to locate, with only a few fragments remaining. These chondrites belong to type H6.

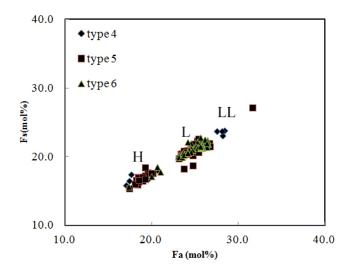


Figure 3. Fa of olivine and Fs of low-Ca pyroxene of 184 chondrites.

3.2. Group L

One hundred and thirty-three meteorites belong to the L group, and they are equilibrated (Table A1). The Fa values of olivine in all the chondrites range from 23.0 to 26.7 mol %, and the Fs values of low-Ca pyroxene range from 18.2 to 22.8 mol %. Both Fa and Fs values are within the ranges of the group L (Figure 3). These chondrites (GRV 020038, 020049, 021643, 021654, 021668, 021726, 021800, and 022221) have a homogenized mineral chemistry and clear outlines of chondrules. Additionally, their fine-grained matrices are only partly recrystallized (Figure 2b), with the sizes of the recrystallized matrices being 5–10 μ m. The primary glass was found in a few chondrules. This indicates that these meteorites experienced the lowest degree of thermal metamorphism. However, all eight chondrites can be assigned to type L4. In 75 meteorites (Table A1), many chondrules are fractured and occur as fragments, but they can be recognized. The matrices of these chondrites are well recrystallized (20–40 μ m), with plagioclase being commonly observed. These meteorites belong to L5. The other 50 meteorites (Table A1) were assigned to type L6. Additionally, their matrices were intensely recrystallized, with the sizes of these grains being 50–200 μ m. Only a few fragments of chondrules can be found (Figure 2c).

3.3. Group LL

Only five chondrites were assigned to this group (Table A1), i.e., GRV 020019, 020021, 020028, 020037, and 020041. The modal abundances of the opaque mineral (0.6–4.1 vol %), along with the compositions of olivine (mean Fa: 27.5–31.6 mol %) and low-Ca pyroxene (mean Fs: 23.1–27.2 mol %), indicate that these meteorites belong to the LL group. GRV 020019 experienced a much stronger thermal metamorphism than the other four meteorites. In GRV 020019, only a few chondrules could be readily recognized. In contrast, in GRV 020021, 020028, 020037, and 020041, the chondrules are clearly outlined (Figure 2d), with their matrices being only partly recrystallized (Figure 2d), and with the sizes of these matrices being 5–20 μ m. Brown-colored glass was found in a few chondrules. Therefore, GRV 020021,020028, 020037, and 020041 were classified as LL4, and GRV 020019 as LL5.

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4. Discussion

4.1. Type and Mass Distribution Patterns

Based on the statistics of the classified 184 equilibrated ordinary chondrites, the relative abundances of H, L, and LL groups were 25%, 72.3%, and 2.7%, respectively. The relative abundance of L chondrites was as high as by a factor of ~2.9 of the H group, while the LL chondrites were rather few in number. Figure 4 shows the type distribution patterns of the Grove Mountain meteorites in this paper when compared with those of Transcontinental Ridge meteorites [16]. The abundance ratio of H:L:LL was 25:72.3:2.7 for meteorites from the Grove Mountains, which is different from that of the Transcontinental Ridges (42:48:10) [16]. It is clear that the Grove Mountain chondrites have a higher abundance of the L chemical group and less H and LL groups when compared with the Transcontinental Ridge chondrites.

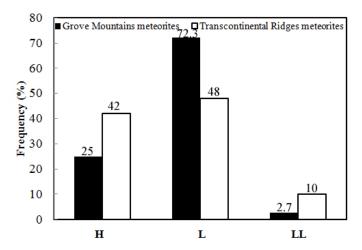


Figure 4. Comparison of the distribution patterns of the chondrite chemical groups between the Transcontinental regions and the Grove Mountains.

The Grove Mountain meteorites have smaller sizes (average mass of 67.9 g) when compared with the Transcontinental Ridge meteorites (average mass of 261.3 g) [16]. The abundance of Grove Mountain meteorites decreases exponentially with the increase of mass. Tiny meteorites (mass < 25 g) are more than 50%. Many Grove Mountain meteorites are small fragments that were found very close to one another. Therefore, they are probably broken pieces from some larger meteorites. On the other hand, the field team in the Grove Mountains searched for meteorites by foot, finding meteorites as small as 0.1 g, while blue ices along the Transcontinental Ridges were swept with a skidoo or snowmobile.

Because of the transfer and concentration of meteorites by glacier and wind, the meteorites collected from the same region in Antarctica face two possibilities: (1) each piece belongs to individual fall events, or (2) some of them may be pieces of the same meteorite fall (meteorite shower, or paired meteorites) [5,16]. Different search methods were employed between the Grove Mountains (by foot) and the Transcontinental Ridges (skidoo), so smaller meteorites could be found in the Grove Mountains. As discussed above, the abundance ratio of L:H was ~2.9; thus, it is obvious that the Grove Mountains have a greater L group presence and a lower H group presence when compared with the Transcontinental Ridges. The abundance ratio of L:H was ~5.6 (Table A1) if we limit ourselves to Grove Mountains meteorites with a mass >10 g. Therefore, the higher H:L ratio in the Grove Mountains is not related to the mass. Both H and L group meteorites were affected, whether they were from a meteor shower or pieces broken from some larger meteorites. In this paper, we reported the petrography and mineral chemistry of 184 equilibrated ordinary chondrites. However, further studies such as terrestrial ages and thermal luminescence are required in order to confirm the pairing.

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4.2. Shock Metamorphism

A petrographic classification of shock metamorphism degrees of ordinary chondrites was put forward by Stöffler et al. [13] and has since been widely used [16–19]. According to their classification, seven shock metamorphism degrees (S1-S6 and shock melted) were defined, based on the shock effects in the main silicate components, as recognized by thin section microscopy. The characteristic shock effects of each shock degree are (Table 1): S1 (unshocked)—sharp optical extinction of olivine; S2 (very weakly shocked)—undulatory extinction of olivine; S3 (weakly shocked)—planar fractures in olivine; S4 (moderately shocked)—mosaicism in olivine; S5 (strongly shocked)—isotropization of plagioclase (maskelynite) and planar deformation features in olivine; and S6 (very strongly shocked—recrystallization of olivine, sometimes combined with phase transformations (ringwoodite and/or phases produced by dissociation reactions). S6 effects are always restricted to regions adjacent to melted portions of a sample which is otherwise only strongly shocked [13]. With the development of high-pressure experiments and study in shock metamorphism in meteorites, modifications and corrections were made on the Pressure-Temperature-time (P-T-t) conditions and formation history of high-pressure minerals in meteorites [11,20–22]. Accordingly, modifications and revisions are needed for the current classification and pressure calibration system of shock degrees in ordinary chondrites (Table 1). In this study, we classified the shock degrees of 184 Grove Mountain ordinary chondrites (see Tables 1 and A1). The shock metamorphism degrees of the 184 chondrites are S1 (10 meteorites), S2 (61 meteorites), S3 (54 meteorites), S4 (54 meteorites), and S5 (five meteorites) (Figure 5), respectively.

Table 1. Shock metamorphism degrees of Grove Mountains chondrites (based on the literature [13]).

Shock Degree		Silicates		I 1F((,)	D It.	
Shock Degree	Olivine	Pyroxene	Plagioclase	Local Effects	Results	
S1	sharp opt	ical extinction a fractures	ınd irregular	none	10	
S2	undulose ex	tinction and irre	egular fractures	none	61	
S3	1	tures undulose ractures (a few		opaque shock veins, incipient formation of melt pockets	54	
S4		fractures weak i gwoodite maske		opaque shock veins and melt pockets interconnecting	54	
S5	ringwo	ractures strong oodite maskelyr Mg, Fe)SiO ₃ -gla	nite (a few	pervasive melt pockets and shock veins	5	
S6		er the shock ind erecrystallization		pervasive melt pockets and shock veins	0	

Figure 6 shows the statistic results of the shock metamorphism degrees within different chemical groups and petrologic types of these Grove Mountain chondrites. Due to its low abundance, the LL group was not considered. The H and L groups displayed distinct abundance patterns. Specifically, 41% of L chondrites were heavily shocked (S4–5) with an occurrence of shock-induced melt veins, while the shock metamorphism is not so extensively developed in H group chondrites. Most H chondrites were weakly shocked, with only four meteorites classified as S4 (Figure 6a).

Stöffler et al. studied the shock metamorphism in 35 H group, 27 L group, and 14 LL group chondrites. Their results showed that the differences in the frequency distribution of shock degrees in H, L, and LL groups are minor; besides, the shock effects and the sequence of progressively increasing degrees of shock metamorphism are very similar [13]. However, the investigation of the Meteoritical Bulletin Database suggests that the frequency distributions of shock degrees between L and H groups are different [11]. There are 23 L group out of all the chondrites that are shocked to S4–S6, whereas the number of H groups is nine. It is evident that the L group chondrites have experienced a stronger

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shock metamorphism than those of the H group [23], which is consistent with the results in this study (Figure 6a).

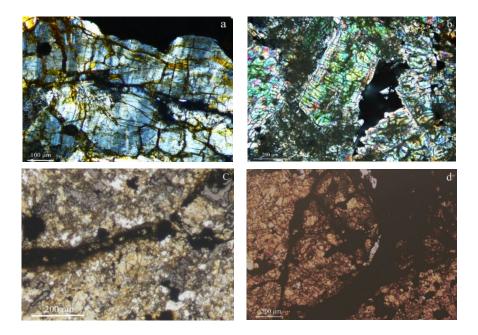


Figure 5. (a) Photomicrograph of olivine in GRV 020069 (S2) under cross-polarized light. The grey-colored grain in the center of the view exhibits undulatory extinction. (b) Photomicrograph of olivine in GRV 022244 (S3) under cross-polarized light. The grain of olivine has two sets of planar fractures. (c) Photomicrograph of shock-induced vein in GRV 020125(S4) under plane-polarized light. Some silicate clasts can be found in vein. (d) Photomicrograph of shock-induced vein in GRV 022194(S5) under plane-polarized light. Melt pockets and shock veins are pervasive.

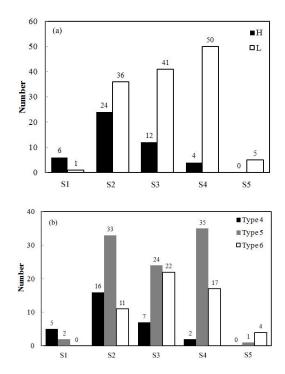


Figure 6. Histograms of shock degrees of Grove Mountain chondrites: (a) chemical groups, and (b) petrographic types.

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Since different groups of chondrites come from different parent bodies, the diverse distribution of frequencies of shock degrees between H and L groups may result from the different physical properties of the surfaces of their parent bodies [24]; for example, the thickness of the soil layer, the petrologic rock type on the surface, etc. As discussed above, a thick and porous regolith is not favorable to reserving high pressure and temperature conditions. An analogous case can be seen on the moon and on Mars. The former is covered by a layer of lunar regolith 2–10 m in depth, while Martian soil is not so abundant on the surface of Mars. Therefore, Martian meteorites have experienced a stronger shock metamorphism than the lunar meteorites [11,25–32].

Another aspect is that the frequency of shock degrees within different petrologic types (Figure 6b) reveals some variations. With increasing petrologic types, the frequency of the shock degrees S3, S4, and S5 increases. The shock degree S2 is the most abundant in type 4 meteorites, while the shock degrees S2 and S4 are the most abundant in type 5 meteorites, and S3 and S4 are the most abundant in type 6 meteorites. There is a lack of shock degrees S1 and S5 in petrologic type 6 and 4 meteorites, respectively. Stöffler et al. made the same observation: no S5 and S6 degrees were observed among type 4 ordinary chondrites [23]. The tendency that, between petrologic types and shock degrees, heavily shocked samples are most likely to be observed in highly petrologic chondrite types may be attributed to the different physical properties of all chondrite types in their unshocked state. High equilibrated samples (type 5/6 chondrites) are essentially coherent nonporous rocks, in which the high pressures and temperatures are more easily retained for a longer time during a shock event, while type 4 chondrites are more porous and richer in volatiles which are more easily crushed upon impact, thereby releasing the high pressure.

5. Conclusions

We studied the petrography and mineral chemistry of 184 Grove Mountain equilibrated ordinary chondrites. Their chemical-petrographic types were assigned, including 46 H groups (22 H4, 20 H5, and four H6), 133 L groups (eight L4, 75 L5, and 50 L6), and five LL groups (four LL4 and one LL5). Some of these chondrites could be paired; however, both H and L group meteorites were affected.

The relative abundances of H, L, and LL are different for Grove Mountain meteorites when compared with those of Transcontinental Ridge meteorites. The Grove Mountain meteorites have smaller sizes (average mass of 67.9 g) in comparison with the latter meteorites (average mass of 261.3 g).

On the basis of the observed shock effects, the shock degrees of the 184 ordinary chondrites were classified. There are 10 with an S1 shock degree, 61 with S2, 54 with S3, 54 with S4, and five with S5. A comparison of the frequencies of shock degrees within different chemical groups shows a diverse frequency distribution of shock degrees between H and L groups, which may represent the difference between the physical properties of their parent body surfaces. Besides, it appears that the frequency of the shock degrees S3, S4, and S5 increases with increasing petrologic types.

Author Contributions: D.D. conceived and designed the experiments; D.D. and S.L. performed the experiments; D.D., S.L. and X.L. analyzed the data; D.D. wrote the paper, assisted by all other authors.

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Conflicts of Interest: The authors declare no conflict of interest.

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Appendix A

Table A1. The major characteristics of 184 ordinary chondrites.

Meteorites	Found Location	Type	Weight (g)	Opaque Mineral (vol%)	Olivine		Low-Ca Pyroxene		- Shock Degrees
ivieteorites					Fa	PMD	Fs	PMD	JHOCK Degrees
GRV020007	S73°05′23′′ E75°11′22′′	H4	25.9	8.4	17.0	3.8	15.8	3.7	S2
GRV020008	S73°05′05′′ E75°12′33′′	H4	6.2	4.5	18.0	1.1	16.1	1.6	S2
GRV020042	S73°00′23′′ E75°11′32′′	H4	2.2	4.2	18.2	1.6	16.2	2.6	S1
GRV020051	S72°59′30′′ E75°13′40′′	H4	1.4	4.0	19.5	1.2	17.4	2.1	S2
GRV020066	S72°59′57′′ E75°14′20′′	H4	1.8	3.3	17.4	4.1	16.5	4.7	S1
GRV020070	S72°59′57′′ E75°12′10′′	H4	7.9	6.0	18.4	0.8	16.4	1.6	S3
GRV020074	S72°59′45′′ E75°12′35′′	H4	1.5	8.4	19.9	6.4	17.5	4.0	S3
GRV020088	S73°00′13′′ E75°15′10′′	H4	3.0	5.1	18.6	1.7	17.0	1.3	S2
GRV020091	S72°59′45′′ E75°13′39′′	H4	5.3	6.1	18.2	1.0	16.4	0.6	S2
GRV020108	S72°58′49′′ E75°15′20′′	H4	5.8	1.4	18.5	1.0	16.5	1.6	S1
GRV020109	S72°58′49′′ E75°15′17′′	H4	2.0	5.4	18.6	0.6	16.5	1.6	S2
GRV020110	S72°58′53′′ E75°14′59′′	H4	3.1	7.3	18.7	1.5	16.6	1.1	S2
GRV020113	S72°58′48′′ E75°15′36′′	H4	1.3	4.8	19.5	8.6	17.5	5.4	S2
GRV020130	S72°58′36′′ E75°15′37′′	H4	139.7	3.3	18.6	0.8	16.6	1.7	S1
GRV021492	S72°58′01′′ E75°16′24′′	H4	211.2	10.0	18.1	1.3	16.0	0.7	S2
GRV021508	S72°56′47′′ E75°17′00′′	H4	278.2	7.4	20.6	1.2	18.1	6.2	S3
GRV021549	S72°56′05′′ E75°19′18′′	H4	16.3	1.5	17.6	1.4	17.4	3.2	S2
GRV021550	S72°56′05′′ E75°19′13′′	H4	18.6	4.3	19.5	4.2	16.7	2.0	S2
GRV021566	S72°56′18′′ E75°17′23′′	H4	13.5	0.8	19.0	2.2	17.3	2.6	S2
GRV021569	S72°56′25′′ E75°17′09′′	H4	19.7	1.7	19.9	2.2	17.5	1.9	S2
GRV021576	S72°57′50′′ E75°12′58′′	H4	105.9	3.4	18.7	1.1	16.9	0.9	S1
GRV021593	S72°51′16′′ E75°12′39′′	H4	28.1	6.0	19.1	1.6	16.9	1.8	S4
GRV020023	S73°05′09′′ E75°12′30′′	H5	2.7	6.1	18.4	0.9	17.0	3.2	S3
GRV020071	S72°59′56′′ E75°12′15′′	H5	4.5	2.7	18.2	1.7	16.6	0.8	S2
GRV020076	S72°59′43′′ E75°12′16′′	H5	37.3	4.5	18.9	3.9	16.6	3.8	S3
GRV020077	S72°59′40′′ E75°12′18′′	H5	1.3	3.2	20.0	3.6	17.6	2.3	S3
GRV020087	S73°00′08′′ E75°16′59′′	H5	2.1	2.4	18.4	1.8	16.6	1.4	S2
GRV020089	S73°00′13′′ E75°15′05′′	H5	4.6	8.5	18.7	1.3	16.9	1.1	S2
GRV020092	S72°59′45′′ E75°12′24′′	H5	1.3	5.0	18.4	2.3	16.9	0.6	S4
GRV020097	S72°59′40′′ E75°12′15′′	H5	131.2	5.5	19.5	2.8	17.7	7.8	S2
GRV020123	S72°58′40′′ E75°15′36′′	H5	3.4	6.2	19.2	1.2	17.2	1.3	S1
GRV020153	S72°58′50′′ E75°15′51′′	H5	1.6	3.4	20.1	1.6	17.6	1.6	S2
GRV020174	S72°58′43′′ E75°15′37′′	H5	1.3	3.7	19.9	1.5	17.6	3.4	S3
GRV021480	S72°56′07′′ E75°19′29′′	H5	12.9	4.3	18.3	3.6	15.9	3.8	S3
GRV021517	S72°56′10′′ E75°18′49′′	H5	96.4	4.0	18.9	1.3	17.0	2.0	S2

Table A1. Cont.

Metaoritas	Found Location	Type	Weight (g)	Opaque Mineral (vol%)	Ol	ivine	Low-Ca Pyroxene		- Shock Degrees
Meteorites	Found Location				Fa	PMD	Fs	PMD	- Shock Degrees
GRV021518	S72°56′11′′ E75°18′51′′	H5	54.1	5.4	18.9	0.8	16.5	2.2	S2
GRV021564	S72°56′03′′ E75°19′20′′	H5	13.7	5.0	18.4	3.6	16.3	2.7	S2
GRV021589	S72°57′57′′ E75°13′48′′	H5	67.4	7.0	19.2	3.2	18.4	6.3	S3
GRV021590	S72°56′08′′ E75°17′54′′	H5	16.5	3.9	19.2	1.2	16.8	2.3	S3
GRV021611	S72°46′31′′ E75°19′29′′	H5	26.7	3.3	18.0	1.1	16.0	0.5	S2
GRV021715	S72°47′24′′ E75°17′47′′	H5	1.5	1.2	17.4	3.9	15.4	1.9	S4
GRV021795	S72°46′26′′ E75°19′26′′	H5	11.8	3.5	18.5	0.7	16.6	0.9	S2
GRV020072	S72°59′54′′ E75°11′45′′	H6	13.0	3.7	20.6	1.9	18.5	5.2	S3
GRV020093	S72°59′44′′ E75°12′23′′	H6	1.1	9.0	19.9	1.5	17.1	1.5	S3
GRV020098	S72°59′40′′ E75°12′16′′	H6	122.7	2.9	21.0	8.9	17.8	3.4	S2
GRV021522	S72°57′34″ E75°14′13″	H6	1.2	6.6	17.3	1.7	15.6	2.5	S4
GRV020038	S73°08′37′′ E75°03′27′′	L4	25.6	3.1	25.7	0.8	22.0	0.8	S2
GRV020049	S72°59′40′′ E75°12′34′′	L4	1.2	4.2	25.1	2.1	21.2	1.4	S2
GRV021643	S72°46′22″ E75°20′44″	L4	12.8	4.7	24.0	1.3	20.4	1.3	S2
GRV021654	S72°46′24″ E75°19′55″	L4	16.6	3.0	25.6	1.8	21.9	2.9	S3
GRV021668	S72°46′43″ E75°19′48″	L4	88.5	3.2	25.1	1.1	21.6	2.6	S3
GRV021726	S72°47′24″ E75°17′22″	L4	11.2	4.0	25.3	1.2	21.5	3.0	S3
GRV021800	S72°46′24″ E75°20′02″	L4	14.1	1.2	25.2	1.7	21.6	2.9	S3
GRV022221	S72°46′42″ E75°19′15″	L4	29.1	3.3	25.0	0.9	21.1	1.5	S4
GRV020040	S73°08′05′′ E75°02′34′′	L5	31.7	6.2	24.4	1.4	21.0	1.9	S2
GRV020068	S72°59′42″ E75°12′21″	L5	623.8	3.5	24.6	0.6	21.2	1.2	S4
GRV020069	S72°59′44″ E75°12′28″	L5	171.5	5.1	24.4	1.5	20.8	0.7	S2
GRV020096	S72°59′49′′ E75°12′22′′	L5	1.8	0.9	24.7	4.9	20.2	20.2	S3
GRV020107	S72°58′48″ E75°15′23″	L5	3.7	1.8	24.2	0.9	20.7	0.8	S4
GRV020125	S72°58′46″ E75°15′37″	L5	2.1	3.9	24.9	0.6	21.4	1.1	S4
GRV020127	S72°58′50′′ E75°15′45′′	L5	6.8	1.6	24.7	1.9	18.7	1.3	S1
GRV020154	S72°58′51″ E75°15′40″	L5	1.5	2.9	26.6	2.9	22.0	1.0	S2
GRV021495	S72°56′48″ E75°18′21″	L5	43.7	3.9	24.2	1.1	20.8	1.1	S4
GRV021499	S72°56′35′′ E75°17′48′′	L5	57.1	3.9	23.1	0.7	19.7	0.7	S2
GRV021500	S72°56′57″ E75°17′33″	L5	12.8	3.9	23.7	1.1	20.8	3.0	S4
GRV021501	S72°56′40′′ E75°17′55′′	L5	17.4	2.5	24.0	1.5	20.7	0.8	S3
GRV021548	S72°57′46′′ E75°14′46′′ S72°57′47′′	L5	31.8	2.4	24.1	0.7	20.9	1.3	S2
GRV021582	E75°13′28′′ S72°57′43′′	L5	14.2	5.3	24.2	2.7	20.5	1.3	S4
GRV021586	E75°15′20′′ S72°57′50′′	L5	11.6	3.4	24.4	1.5	21.0	1.1	S2
GRV021587	E75°13′31′′ S72°46′24′′	L5	31.3	3.7	24.2	2.7	20.5	1.0	S2
GRV021614	E75°19′56′′ S72°46′35′′	L5	17.3	3.1	23.4	1.1	20.2	1.0	S4
GRV021651	E75°20′52′′	L5	35.2	4.1	24.6	0.7	21.1	0.8	S3
GRV021652	S72°46′28′′ E75°20′03′′	L5	57.9	2.8	25.8	2.2	21.6	2.9	S3

Table A1. Cont.

Meteorites	Found Location	Type	e Weight (g)	Opaque	Olivine		Low-Ca Pyroxene		- Shock Degrees
		туре	vveigitt (g)	Mineral (vol%)	Fa	PMD	Fs	PMD	- Shock Degrees
GRV021669	S72°46′42′′ E75°19′49′′	L5	406.1	2.4	26.0	2.6	22.1	2.5	S3
GRV021670	S72°46′43′′ E75°19′54′′	L5	282.4	3.0	26.0	2.3	21.9	2.6	S3
GRV021724	S72°47′23′′ E75°17′50′′	L5	22.0	3.3	23.4	1.4	20.2	0.9	S4
GRV021725	S72°47′23′′ E75°17′31′′	L5	25.0	3.3	26.1	2.9	21.8	3.4	S2
GRV021785	S72°46′30′′ E75°20′20′′	L5	71.7	2.9	25.6	0.9	21.6	3.2	S3
GRV021794	S72°46′25′′ E75°19′23′′	L5	44.5	0.2	24.2	1.9	20.9	0.5	S3
GRV021801	S72°46′24′′ E75°20′03′′	L5	28.1	1.6	24.9	1.9	21.8	2.3	S2
GRV021802	S72°46′24′′ E75°20′04′′	L5	175.0	2.4	25.7	1.0	21.6	2.2	S2
GRV021803	S72°46′27′′ E75°20′20′′	L5	115.9	5.7	25.4	1.9	21.4	3.3	S4
GRV021804	S72°46′30′′ E75°19′46′′	L5	46.1	1.7	25.6	1.8	21.5	2.9	S4
GRV022024	S72°47′01′′ E75°17′58′′	L5	105.9	1.8	23.4	2.1	20.1	0.8	S2
GRV022026	S72°47′00′′ E75°17′48′′	L5	73.3	3.3	23.5	1.4	20.2	2.0	S2
GRV022027	S72°46′59′′ E75°17′54′′	L5	169.7	3.3	23.4	1.2	20.1	1.1	S4
GRV022028	S72°46′57′′ E75°18′01′′	L5	244.9	2.1	23.9	1.6	20.4	1.7	S2
GRV022038	S72°46′54′′ E75°18′10′′	L5	1012.0	3.5	25.2	1.3	22.2	3.6	S3
GRV022039	S72°46′43′′ E75°18′55′′	L5	160.7	2.0	24.8	1.0	20.9	1.8	S3
GRV022040	S72°46′43′′ E75°18′55′′	L5	589.5	2.8	24.8	2.0	21.5	3.9	S3
GRV022041	S72°46′43′′ E75°18′55′′	L5	96.4	2.2	25.6	1.5	21.5	1.6	S4
GRV022042	S72°46′43′′ E75°18′55′′	L5	218.1	2.5	26.7	2.9	21.5	3.4	S2
GRV022126	S72°46′44′′ E75°18′49′′	L5	48.0	3.1	23.5	1.0	20.5	2.6	S2
GRV022127	S72°46′44′′ E75°18′49′′	L5	39.2	2.4	23.9	1.3	20.7	2.6	S4
GRV022128	S72°46′44′′ E75°18′49′′	L5	31.1	2.8	23.6	1.8	20.4	0.9	S4
GRV022141	S72°46′44′′ E75°18′49′′	L5	26.6	2.3	23.6	2.4	20.5	2.1	S4
GRV022142	S72°46′44′′ E75°18′49′′	L5	23.3	4.9	24.0	3.5	20.8	2.6	S4
GRV022143	S72°46′44′′ E75°18′49′′	L5	14.7	3.4	23.4	1.6	20.1	0.7	S4
GRV022146	S72°46′44′′ E75°18′49′′	L5	20.7	2.6	23.7	2.8	20.4	1.0	S4
GRV022150	S72°46′44′′ E75°18′49′′	L5	14.3	4.2	23.7	12.4	18.2	4.4	S4
GRV022151	S72°46′44′′ E75°18′49′′	L5	12.5	3.8	25.2	1.4	22.3	2.4	S4
GRV022159	S72°46′54″ E75°18′14″	L5	13.0	0.2	23.4	2.7	20.0	0.8	S4
GRV022160	S72°46′54′′ E75°18′14′′	L5	610.1	3.2	23.7	1.8	20.4	4.6	S4
GRV022161	S72°46′54″ E75°18′14″	L5	26.5	5.8	24.0	1.8	20.6	1.0	S4
GRV022163	S72°46′54′′ E75°18′14′′	L5	94.3	3.8	25.7	1.8	21.4	2.3	S3
GRV022168	S72°46′54′′	L5	139.7	2.1	25.2	0.7	21.2	1.4	S4
GRV022169	E75°18′14″ S72°46′54″	L5	93.1	3.6	25.4	3.3	21.5	4.7	S4
GRV022170	E75°18′10″ S72°46′54″	L5	80.0	3.6	25.4	2.0	22.6	2.0	S5
GRV022177	E75°18′10″ S72°46′54″	L5	34.5	3.1	23.6	1.5	20.2	1.1	S4
GRV022185	E75°18′10″ S72°46′52″	L5	13.3	3.7	23.8	1.6	20.5	1.2	S2
GRV022186	E75°17′32′′ S72°46′52′′	L5	12.0	4.9	23.5	1.1	20.4	0.8	S2
GRV022190	E75°17′32′′ S72°46′53′′	L5	42.1	6.0	23.6	0.8	20.4	1.3	S2
GAV 022190	E75°18′20′′	LO	44.1	0.0	23.0	0.0	20.2	1.3	54

Table A1. Cont.

Meteorites	Found Location	Type	e Weight (g)	Opaque Mineral (vol%)	Olivine		Low-Ca	Pyroxene	- Shock Degrees	
		-71-			Fa	PMD	Fs	PMD		
GRV022191	S72°46′53′′ E75°18′20′′	L5	18.8	2.6	26.0	2.6	21.7	2.6	S4	
GRV022192	S72°46′53′′ E75°18′20′′	L5	19.0	3.2	25.4	1.9	20.7	1.3	S4	
GRV022193	S72°46′53′′ E75°18′20′′	L5	14.8	3.9	25.4	1.7	21.9	2.6	S4	
GRV022199	S72°46′51′′ E75°18′08′′	L5	30.1	3.6	25.6	3.2	22.1	2.6	S3	
GRV022206	S72°46′51′′ E75°18′08′′	L5	22.5	4.7	23.7	1.6	20.8	0.8	S2	
GRV022207	S72°46′51′′ E75°18′08′′	L5	15.8	4.7	23.4	0.8	20.0	1.5	S4	
GRV022210	S72°46′42′′ E75°19′15′′	L5	11.8	2.0	23.3	1.0	20.4	1.1	S4	
GRV022211	S72°46′42′′ E75°19′15′′	L5	11.1	3.7	23.5	1.0	20.2	1.2	S2	
GRV022212	S72°46′42′′ E75°19′15′′	L5	12.1	2.2	23.9	3.5	20.4	2.7	S2	
GRV022219	S72°46′42′′ E75°19′15′′	L5	101.0	3.6	25.6	1.8	22.0	5.9	S4	
GRV022222	S72°46′42′′ E75°19′15′′	L5	126.0	2.2	25.5	2.0	21.6	2.1	S3	
GRV022284	S72°46′47′′ E75°17′34′′	L5	11.9	3.2	24.0	1.2	20.7	1.3	S2	
GRV022285	S72°46′47′′ E75°17′34′′	L5	37.1	4.1	23.8	0.8	20.4	0.9	S2	
GRV022287	S72°46′47′′ E75°17′34′′	L5	11.1	3.1	23.4	2.9	20.2	1.1	S4	
GRV022291	S72°46′46′′ E75°17′43′′	L5	12.2	2.8	25.1	1.5	21.8	6.4	S3	
GRV022443	S72°46′23′′ E75°20′21′′	L5	720.9	3.2	24.9	1.5	21.9	3.6	S3	
GRV022444	S72°46′23′′ E75°20′21′′	L5	679.2	3.3	25.4	1.4	21.4	3.5	S3	
GRV020027	S73°04′09′′ E75°16′37′′	L6	428.8	3.8	25.8	1.6	21.9	1.8	S4	
GRV020073	S72°59′56′′ E75°12′46′′	L6	9.0	1.8	26.4	2.8	22.4	1.2	S3	
GRV020095	S72°59′43′′ E75°12′24′′	L6	6.1	1.4	26.5	1.3	21.9	1.6	S5	
GRV020115	S72°58′43′′ E75°15′58′′	L6	8.3	1.1	25.8	1.3	22.0	1.9	S2	
GRV020131	S72°58′37′′ E75°15′40′′	L6	3.8	3.9	26.0	4.1	21.8	1.1	S2	
GRV020132	S72°58′38′′ E75°15′41′′	L6	1.9	3.4	25.8	3.3	21.7	2.2	S2	
GRV020134	S72°58′41′′ E75°15′38′′	L6	16.4	3.8	26.4	3.0	22.2	2.3	S2	
GRV020163	S72°58′35′′ E75°15′54′′	L6	3.4	5.4	24.7	0.8	21.0	0.8	S4	
GRV020167	S72°58′41′′ E75°15′49′′	L6	7.5	4.1	25.9	2.3	21.7	1.1	S3	
GRV021475	S72°56′35′′ E75°16′30′′	L6	183.7	2.5	25.2	2.5	22.1	2.5	S2	
GRV021478	S72°56′06′′ E75°19′30′′	L6	104.5	4.4	26.4	2.2	21.9	1.0	S2	
GRV021502	S72°57′06′′ E75°16′25′′	L6	17.4	2.2	26.0	1.4	21.9	3.0	S3	
GRV021503	S72°56′54′′ E75°15′56′′	L6	23.6	3.3	26.1	0.8	21.8	1.0	S3	
GRV021504	S72°56′55′′ E75°15′54′′	L6	17.8	1.5	26.1	1.8	21.8	0.7	S3	
GRV021506	S72°56′42′′ E75°15′39′′	L6	39.1	3.5	26.1	2.0	22.5	3.9	S3	
GRV021578	S72°57′46′′ E75°14′20′′	L6	14.4	4.0	24.3	0.8	21.0	1.0	S2	
GRV021595	S72°49′28′′ E75°17′30′′	L6	22.5	2.7	26.3	2.5	21.8	2.2	S3	
GRV021597	S72°49′29′′ E75°17′29′′	L6	26.8	4.3	25.8	2.1	21.7	1.2	S3	
GRV021649	S72°46′24′′ E75°20′00′′	L6	13.9	2.5	23.0	2.8	20.1	1.4	S5	
GRV021714	S72°47′24′′ E75°17′50′′	L6	136.8	2.9	23.4	1.8	20.2	1.6	S4	
GRV021722	S72°47′26′′ E75°16′58′′	L6	33.2	3.1	23.7	1.4	20.4	1.0	S4	
GRV021723	S72°47′26′′ E75°16′37′′	L6	25.6	2.4	23.3	1.6	20.1	1.8	S4	
GRV021786	S72°46′31′′ E75°20′20′′	L6	72.5	4.0	25.1	2.3	21.5	1.6	S3	

Table A1. Cont.

Meteorites	Found Location	Type	Weight (g)	Opaque Mineral (vol%)	Olivine		Low-Ca Pyroxene		- Shock Degrees
	Found Location	туре			Fa	PMD	Fs	PMD	- Shock Degrees
GRV021787	S72°46′30′′ E75°20′30′′	L6	70.1	4.1	26.3	2.3	21.4	2.7	S3
GRV021796	S72°46′26′′ E75°19′27′′	L6	55.3	1.0	23.7	1.6	20.4	1.7	S4
GRV021797	S72°46′26′′ E75°19′28′′	L6	99.7	3.3	23.5	1.1	20.2	1.1	S3
GRV021799	S72°46′24′′ E75°20′01′′	L6	49.8	1.5	23.2	1.3	20.1	1.7	S4
GRV022025	S72°47′01′′ E75°17′47′′	L6	101.4	1.3	23.5	2.8	20.2	1.5	S2
GRV022114	S72°46′45′′ E75°18′38′′	L6	171.4	1.9	23.1	1.6	19.9	1.4	S4
GRV022129	S72°46′44′′ E75°18′49′′	L6	30.2	1.8	23.6	0.9	20.4	0.7	S4
GRV022130	S72°46′44′′ E75°18′49′′	L6	30.5	2.0	24.5	2.9	22.0	2.8	S3
GRV022131	S72°46′44′′ E75°18′49′′	L6	24.8	3.3	25.5	1.8	21.8	1.8	S3
GRV022132	S72°46′44″ E75°18′49″	L6	23.9	3.2	25.4	1.9	21.3	1.7	S3
GRV022133	S72°46′44′′ E75°18′49′′	L6	21.4	3.2	25.2	0.4	21.3	1.3	S3
GRV022134	S72°46′44′′ E75°18′49′′	L6	20.0	1.1	25.7	2.5	21.5	4.0	S3
GRV022145	S72°46′44′′ E75°18′49′′	L6	31.8	3.3	23.5	1.0	19.9	0.9	S2
GRV022147	S72°46′44′′ E75°18′49′′	L6	21.2	3.0	25.7	1.2	21.2	1.7	S4
GRV022148	S72°46′44′′ E75°18′49′′	L6	17.2	2.3	25.3	1.6	21.3	2.3	S4
GRV022149	S72°46′44′′ E75°18′49′′	L6	17.2	3.6	25.9	1.2	21.3	2.2	S4
GRV022158	S72°46′44″ E75°18′49″	L6	14.2	1.5	23.6	2.0	20.6	4.4	S4
GRV022162	S72°46′54″ E75°18′14″	L6	141.0	1.8	24.1	1.7	20.7	1.0	S3
GRV022164	S72°46′54′′ E75°18′14′′	L6	14.9	2.9	25.6	1.8	22.8	3.8	S4
GRV022178	S72°46′54′′ E75°18′10′′	L6	23.9	3.5	23.7	1.4	20.3	1.0	S4
GRV022194	S72°46′53′′ E75°18′20′′	L6	15.6	2.8	25.8	3.3	21.3	3.8	S5
GRV022220	S72°46′42″ E75°19′15″	L6	12.0	2.9	26.3	3.4	22.0	4.8	S3
GRV022223	S72°46′42″ E75°19′15″	L6	71.2	3.1	25.7	1.7	21.4	3.4	S3
GRV022237	S72°46′43″ E75°18′55″	L6	14.1	12.0	24.5	0.4	20.7	0.8	S5
GRV022282	S72°46′42′′ E75°18′50′′	L6	43.3	3.3	24.1	1.3	20.6	1.5	S2
GRV022288	S72°46′47′′ E75°17′34′′	L6	33.0	2.4	25.6	3.1	21.5	3.5	S3
GRV022289	S72°46′47′′ E75°17′34′′	L6	38.1	2.1	24.1	4.9	22.1	3.7	S4
GRV020021	S73°05′11″ E75°13′59″	LL4	11.3	2.2	27.5	3.4	23.7	4.2	S1
GRV020028	S73°04′23″ E75°16′08″	LL4	10.1	3.3	28.2	1.0	23.1	2.5	S3
GRV020037	S73°06′20″ E75°10′14″	LL4	1.6	2.8	28.4	1.2	23.8	4.6	S1
GRV020041	S73°06′19″ E75°09′08″	LL4	1.4	4.1	28.1	0.9	23.7	4.0	S1
GRV020019	S73°05′02′′ E75°12′59′′	LL5	10.1	0.6	31.6	2.3	27.2	2.6	S2

References

- 1. Liu, X.; Zhao, Y.; Liu, X.; Yu, L. Geological characteristics of Grove Mountains, East Antarctica: New proof of final suture of Gondwana. *Chin. Sci. D* **2002**, *32*, 457–468.
- 2. Kimura, M.; Lin, Y.; Ikeda, Y.; Goresy, A.E.; Yanai, K.; Kojima, H. Mineralogy of Antarctic aubrites, Yamato-793592 and Allan Hills-78113: Comparison with non-Antarctic aubrites and E-chondrites. *Antarct. Meteor. Res.* **1993**, *6*, 186–203.

3. Nobuyoshi, T.; Haramura, H.; Ikeda, Y.; Kimura, M.; Kojima, H.; Imae, N.; Lee, M.S. Comparative study of the major element chemical compositions of Antarctic chondrites to those of non-Antarcitc falls with reference to terrestrial weathering. *Antarct. Meteor. Res.* **1997**, *10*, 165–180.

- 4. Takeda, H.; Mori, H.; Saito, J.; Miyamoto, M. Mineral-chemical comparisons of MAC88105 with Yamato lunar meteorites. *Geochim. Cosmochim. Acta* **1991**, *55*, 3009–3017. [CrossRef]
- 5. Welten, K.C.; Nishiizumi, K.; Caffee, M.W.; Hillegonds, D.J.; Johnson, J.A.; Jull, A.J.T.; Wieler, R.; Folco, L. Terrestrial ages, pairing, and concentration mechanism of Antarctic chondrites from Frontier Mountain, Northern Victoria Land. *Meteorit. Planet. Sci.* **2006**, *41*, 1081–1094. [CrossRef]
- Chen, M.; El Goresy, A.; Gillet, P. Ringwoodite lamellae in olivine: Clues to olivine-ringwoodite phase transition mechanisms in shocked meteorites and subducting slabs. *Proc. Natl. Acad. Sci. USA* 2004, 101, 15033–15037.
 [CrossRef] [PubMed]
- 7. Tomioka, N.; Miyahara, M. High-pressure minerals in shocked meteorites. *Meteorit. Planet. Sci.* **2017**, 52, 2017–2039. [CrossRef]
- 8. Miyahara, M.; El Goresy, A.; Ohtani, E.; Nagase, T.; Nishijima, M.; Vashaei, Z.; Ferroir, T.; Gillet, P.; Dubrovinsky, L.; Simionovici, A. Evidence for fractional crystallization of wadsleyite and ringwoodite from olivine melts in chondrules entrained in shock-melt veins. *Proc. Natl. Acad. Sci. USA* **2008**, *105*, 8542–8547. [CrossRef] [PubMed]
- 9. Yin, F.; Liao, Z.; Hursthouse, A.; Dai, D. Shock-Induced Olivine-Ringwoodite Transformation in the Shock Vein of Chondrite GRV053584. *Minerals* **2008**. [CrossRef]
- 10. Feng, L.; Lin, Y.; Hu, S.; Xu, L.; Miao, B. Estimating compositions of natural ringwoodite in the heavily shocked Grove Mountains 052049 meteorite from Raman spectra. *Am. Mineral.* **2011**, *96*, 1480–1489. [CrossRef]
- 11. Feng, L.; Lin, Y.; Hu, S.; Liu, T. Shock metamorphism of ordinary chondrites from Grove Mountains, Antarctica. *Chin. J. Pol. Sci.* **2009**, *20*, 187–199.
- 12. Lin, Y.; Wang, D.; Ouyang, Z. Progress in study of Chinese Antarctica Meteorites. *Chin. J. Pol. Sci.* **2009**, 20, 81–96.
- 13. Stöffler, D.; Keil, K.; Scott, E.R.D. Shock metamorphism of ordinary chondrites. *Geochim. Cosmochim. Acta* **1991**, *55*, 3845–3867. [CrossRef]
- 14. Miao, B. The status of Antarctic meteorite research in China and the prospect on its development. *Bull. Mineral. Petrol. Geochem.* **2015**, *34*, 1081–1089. (In Chinese)
- 15. Van Schmus, W.R.; Wood, J.A. A chemical-petrologic classification for the chondritic meteorites. *Geochim. Cosmochim. Acta* **1967**, *31*, 747–765. [CrossRef]
- 16. Miao, B.; Lin, Y.; Ouyang, Z.; Zhou, X. Type distribution pattern and pairing of ordinary chondrites from Grove Mountains, Antarctica. *Chin. Sci. Bull.* **2003**, *48*, 908–913.
- 17. Miao, B.; Wang, D. Classification of Grove Mountains meteorites and its significance. *Chin. J. Pol. Sci.* **2009**, 20, 97–109.
- 18. Lu, R.; Miao, B.; Wang, G.; Dai, D.; Lin, Y.; Ouyang, Z.; Li, C. Classification of 24 New Ordinary Chondrites from Grove Mountains, Antarctica. *Acta Geol. Sin.* **2004**, *78*, 1052–1059.
- 19. Fritz, J.; Greshake, A.; Fernandes, V.A. Revising the shock classification of meteorites. *Meteorit. Planet. Sci.* **2017**, 52, 1216–1232. [CrossRef]
- 20. Rubie, D.C.; Ross, C.R. Kinetics of the olivine-spinel transformation in subducting lithosphere: Experimental constraints and implications for deep slab processes. *Phys. Earth Planet. Int.* **1994**, *86*, 223–243. [CrossRef]
- 21. Pittarello, L.; Ji, G.; Yamaguchi, A.; Schryvers, D.; Debaille, V.; Claeys, P. From olivine to ringwoodite: A TEM study of a complex process. *Meteorit. Planet. Sci.* **2015**, *50*, 944–957. [CrossRef]
- 22. Chen, M.; Li, H.; El Goresy, A.; Liu, J.; Xie, X. Fracture-related intracrystalline transformation of olivine to ringwoodite in the shocked Sixiangkou meteorite. *Meteorit. Planet. Sci.* **2006**, *41*, 731–737. [CrossRef]
- 23. Stöffler, D.; Keil, K.; Scott, E.R.D. Shock classification of ordinary chondrites: New data and interpretations. *Meteoritics* **1992**, *27*, 292–293.
- 24. Sharp, T.G.; Xie, Z.; DeCarli, P.S.; Hu, J. A large shock vein in L chondrite Roosevelt County 106: Evidence for a long-duration shock pulse on the L chondrite parent body. *Meteorit. Planet. Sci.* **2015**, *50*, 1941–1953. [CrossRef]

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25. Walton, E.L. Shock metamorphism of Elephant Moraine A79001: Implications for olivine–ringwoodite transformation and the complex thermal history of heavily shocked Martian meteorites. *Geochim. Cosmochim. Acta* 2013, 107, 299–315. [CrossRef]

- 26. Greshake, A.; Fritz, J.; Böttger, U.; Goran, D. Shear-induced ringwoodite formation in the Martian shergottite Dar al Gani 670. *Earth Planet. Sci. Lett.* **2013**, 375, 383–394. [CrossRef]
- 27. Ma, C.; Tschauner, O.; Beckett, J.R.; Liu, Y.; Rossman, G.R.; Sinogeikin, S.V.; Smith, J.S.; Taylor, L.A. Ahrensite, γ-Fe₂SiO₄, a new shock-metamorphic mineral from the Tissint meteorite: Implications for the Tissint shock event on Mars. *Geochim. Cosmochim. Acta* **2016**, *184*, 240–256. [CrossRef]
- 28. Barrat, J.A.; Chaussidon, M.; Bohn, M.; Gillet, P.; Göpel, C.; Lesourd, M. Lithium behavior during cooling of a dry basalt: An ion-microprobe study of the lunar meteorite Northwest Africa 479 (NWA 479). *Geochim. Cosmochim. Acta* 2005, 69, 5597–5609. [CrossRef]
- 29. Zhang, A.C.; Hsu, W.B.; Floss, C.; Li, X.H.; Li, Q.L.; Liu, Y.; Taylor, L.A. Petrogenesis of lunar meteorite Northwest Africa 2977: Constrains from in situ microprobe results. *Meteorit. Planet. Sci.* **2010**, 45, 1929–1947. [CrossRef]
- 30. Ohtani, E.; Ozawa, S.; Miyahara, M.; Ito, Y.; Mikouchi, T.; Kimura, M.; Arai, T.; Sato, K.; Hiraga, K. Coesite and stishovite in a shocked lunar meteorite, Asuka-881757, and impact events in lunar surface. *Proc. Natl. Acad. Sci. USA* **2011**, *108*, 463–466. [CrossRef] [PubMed]
- 31. Miyahara, M.; Kaneko, S.; Ohtani, E.; Sakai, T.; Nagase, T.; Kayama, M.; Nishido, H.; Hirao, N. Discovery of seifertite in a shocked lunar meteorite. *Nat. Commun.* **2013**, *4*. [CrossRef] [PubMed]
- 32. Kaneko, S.; Miyahara, M.; Ohtani, E.; Arai, T.; Hirao, N.; Sato, K. Discovery of stishovite in Apollo 15299 sample. *Am. Mineral.* **2015**, *100*, 1308–1311. [CrossRef]



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